

*Reviews of Geophysics*

Supporting Information for

**Climatic changes and their elevational patterns in the mountains of the world**

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## **Contents of this file**

Tables S1 to S3

Text T1 to T4

## **Introduction**

This file contains information as follows

Table S1: List of CMIP5 models followed by references Text T1

Text T2: Details of observational datasets

Text T3: Metadata on SROCC station analyses

Text T4: List of references from SROCC Tables

Tables S2 and S3: Detailed regional trends in temperature and precipitation

<b>Model ID</b>	<b>Resolution Lon × Lat° Lev</b>	<b>Institution ID</b>	<b>Key reference</b>
<b>ACCESS1-0</b>	<b>1.875 × 1.25L38 (N96)</b>	<b>CSIRO-BOM</b>	<b>Bi et al. (2013)</b>
<b>ACCESS1-3</b>	<b>1.875 × 1.25L38</b>	<b>CSIRO-BOM</b>	<b>Bi et al. (2013)</b>
<b>bcc-csm1-1-m</b>	<b>1.125 × 1.125L26 (T106)</b>	<b>BCC</b>	<b>Wu et al. (2013)</b>
<b>bcc-csm1-1</b>	<b>2.8125 × 2.8125L26 (T42)</b>	<b>BCC</b>	<b>Wu et al. (2013)</b>
<b>BNU-ESM</b>	<b>2.8125 × 2.8125L26 (T42)</b>	<b>BNU</b>	<b>Ji et al. (2014)</b>
<b>CanESM2</b>	<b>2.8125 × 2.8125L35 (T63)</b>	<b>CCCMA</b>	<b>Arora et al. (2011)</b>
<b>CCSM4</b>	<b>1.25 × 0.9L27 (T63)</b>	<b>NCAR</b>	<b>Meehl et al. (2012)</b>
<b>CESM1-BGC</b>	<b>1.25 × 0.9L27</b>	<b>NSF-DOE- NCAR</b>	<b>Hurrell et al. (2013)</b>
<b>CESM1-CAM5</b>	<b>1.25 × 0.9L27</b>	<b>NSF-DOE- NCAR</b>	<b>Hurrell et al. (2013)</b>
<b>CESM1- FASTCHEM</b>	<b>1.25 × 0.9L27</b>	<b>NSF-DOE- NCAR</b>	<b>Hurrell et al. (2013)</b>
<b>CESM1-WACCM</b>	<b>2.5 × 1.9</b>	<b>NSF-DOE- NCAR</b>	<b>Hurrell et al. (2013)</b>
<b>CMCC-CESM</b>	<b>3.75 x 3.75</b>	<b>CMCC</b>	<b>Hurrell et al. (2013)</b>
<b>CMCC-CM</b>	<b>0.75 × 0.75L31 (T159)</b>	<b>CMCC</b>	<b>Scoccimarro et al. (2011)</b>
<b>CMCC-CMS</b>	<b>1.875 × 1.875L95 (T63)</b>	<b>CMCC</b>	<b>Davini et al. (2013)</b>
<b>CNRM-CM5</b>	<b>1.40625 × 1.40625L31 (T127)</b>	<b>CNRM- CERFACS</b>	<b>Voldoire et al. (2013)</b>
<b>CNRM-CM5-2</b>	<b>1.40625 × 2.8125</b>	<b>CNRM- CERFACS</b>	<b>Voldoire et al. (2013)</b>
<b>CSIRO-Mk3-6-0</b>	<b>1.875 × 1.875L18 (T63)</b>	<b>CSIRO- QCCCE</b>	<b>Rotstayn et al. (2012)</b>
<b>EC-Earth</b>	<b>1.125 × 1.125L62 (T159)</b>	<b>EC-EARTH</b>	<b>Hazeleger et al. (2012)</b>
<b>FGOALS-g2</b>	<b>2.8125 × 2.8125L26</b>	<b>LASG-CESS</b>	<b>Li et al. (2013)</b>
<b>FIO-ESM</b>	<b>2.8125 × 2.8125L26 (T42)</b>	<b>FIO</b>	<b>Song et al. (2012)</b>

<b>GFDL-ESM2G</b>	<b>2.5 × 2L24 (M45)</b>	<b>GFDL</b>	<b>Delworth et al. (2006)</b>
<b>GFDL-ESM2M</b>	<b>2.5 × 2L24 (M45)</b>	<b>GFDL</b>	<b>Delworth et al. (2006)</b>
<b>GFDL-CM2p1</b>	<b>2.5 × 2L24 (M45)</b>	<b>GFDL</b>	<b>Yang et al. (2020)</b>
<b>GFDL-CM3</b>	<b>2.5 × 2L48 (C48)</b>	<b>GFDL</b>	<b>Delworth et al. (2006)</b>
<b>GISS-E2-H-CC</b>	<b>2×2.5×L40</b>	<b>NASA GISS</b>	<b>Miller et al. (2014)</b>
<b>GISS-E2-R-CC</b>	<b>2×2.5×L40</b>	<b>NASA GISS</b>	<b>Miller et al. (2014)</b>
<b>GISS-E2-R</b>	<b>2×2.5×L40</b>	<b>NASA GISS</b>	<b>Miller et al. (2014)</b>
<b>HadCM3</b>	<b>3.75 x 2.5(L19)</b>	<b>MOHC</b>	<b>Johns et al. (2003)</b>
<b>HadGEM2-AO</b>	<b>1.875 × 1.24L60</b>	<b>MOHC</b>	<b>Martin et al. (2011)</b>
<b>HadGEM2-CC</b>	<b>1.875 × 1.24L60 (N96)</b>	<b>MOHC</b>	<b>Martin et al. (2011)</b>
<b>HadGEM2-ES</b>	<b>1.875 × 1.24L38 (N96)</b>	<b>MOHC</b>	<b>Bellouin et al. (2011)</b>
<b>INM-CM4</b>	<b>2 × 1.5L21</b>	<b>INM</b>	<b>Volodin et al. (2010)</b>
<b>IPSL-CM5A-LR</b>	<b>3.75 × 1.89L39</b>	<b>IPSL</b>	<b>Hourdin et al. (2013)</b>
<b>IPSL-CM5A-MR</b>	<b>2.5 × 1.2587L39</b>	<b>IPSL</b>	<b>Hourdin et al. (2013)</b>
<b>IPSL-CM5B-LR</b>	<b>3.75 × 1.9L39</b>	<b>IPSL</b>	<b>Hourdin et al. (2013)</b>
<b>MIROC5</b>	<b>1.40625 × 1.40625L40 (T85)</b>	<b>MIROC</b>	<b>Watanabe et al. (2010)</b>
<b>MIROC-ESM</b>	<b>2.8125 × 2.8125L80 (T42)</b>	<b>MIROC</b>	<b>Watanabe et al. (2011)</b>
<b>MIROC-ESM-CHEM</b>	<b>2.8125 × 2.8125L80 (T42)</b>	<b>MIROC</b>	<b>Watanabe et al. (2011)</b>
<b>MPI-ESM-LR</b>	<b>1.875 × 1.875L47 (T63)</b>	<b>MPI-M</b>	<b>Giorgetta et al. (2013)</b>
<b>MPI-ESM-MR</b>	<b>1.875 × 1.875L95 (T63)</b>	<b>MPI-M</b>	<b>Giorgetta et al. (2013)</b>
<b>MPI-ESM-P</b>	<b>1.875 × 1.875 (T63)</b>	<b>MPI</b>	<b>Giorgetta et al. (2013)</b>
<b>MRI-CGCM3</b>	<b>1.125 × 1.125L48 (T159)</b>	<b>MRI</b>	<b>Yukimoto et al. (2012)</b>
<b>NorESM1-M</b>	<b>2.5 × 1.9L26 (F19)</b>	<b>NCC</b>	<b>Bentsen et al. (2013)</b>
<b>NorESM1-ME</b>	<b>2.5 × 1.9L26</b>	<b>NCC</b>	<b>Bentsen et al. (2013)</b>

**Table S1: List of CMIP5 models used to create the model ensemble in this study.**

## Text T1: References for Table S1

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## **Text: T2: Descriptions of Gridded Datasets used in this study**

The Climatic Research Unit gridded Time Series (**CRU TS**) dataset (Harris et al., 2020) is a high-resolution (0.5°x0.5°) monthly gridded database of station observations with global coverage (land-only, excluding Antarctica) covering 1901–2018. It provides both near surface temperature (K) and total precipitation (mm/month) based on over 4,000 stations (10,000 since the 1970s). Observations from individual stations are normalized to their 1961–1990 mean, gridded to the regular grid using angular distance weighting, and resulting anomalies converted to absolute values using the CRU CL v1.0 climatology. Synthesized observations (at 2.5° resolution) are used to fill gaps in regions of poor observational coverage.

The GISS Surface Temperature (**GISTEMP**) dataset provides monthly estimates of global surface temperature anomalies covering the time period 1880–2019 over a grid of 2x2° latitude-longitude resolution. The version employed in this study is GISTEMP v4 (with a 250 km smoothing radius, i.e. the distance in kilometers over which a station influences regional temperature (Hansen et al., 2010, Lenssen et al., 2019)).

The Global Precipitation Climatology Centre (**GPCC**) reanalysis is a precipitation gridded dataset based on gauge measurements. It is particularly recommended for model verification and for analysis of historic global precipitation, or for research on the global water cycle (Becker et al. 2013, Schneider et al. 2017). GPCC version 2018 data extend from 1901 to 2018 and consist of monthly precipitation at 0.25°x0.25° spatial resolution. The data coverage per month varies from ~10,000 stations at the beginning of the 20th century to more than 45,000 stations in 1986/1987. For further details about the spatial



distribution of the stations underlying the GPCC dataset see Schneider et al. (2017). The GPCC monthly precipitation anomalies are spatially interpolated using a spherical adaptation of Shepard's empirical weighting scheme. The latest version also includes estimates of the mean systematic gauge-error.

**ERA5** is the latest reanalysis dataset of the European Centre for Medium-Range Weather Forecasts (Hersbach et al., 2020). It largely improves on the previously widely adopted ERA-Interim data in terms of resolution and physics. It merges information from a vast range of observations with models and assimilation techniques into a global  $0.25^{\circ} \times 0.25^{\circ}$  dataset, covering 1979 to present at the time of compiling this article. The newest version includes information on uncertainty through an ensemble approach. A very recent extension back to 1950 was not used.

An ensemble of **Global climate and Earth system models** from the Coupled Model Intercomparison Project Phase 5 (hereafter **CMIP5**) were considered in this study. We adopted temperature and precipitation output of historical (1900–2005) simulations of 32 models from the CMIP5 project (Taylor et al., 2012) available from the Earth System Grid Federation archive data portal (<http://esgf.llnl.gov>). The full list of models is reported in Table S1. The models are based on a variety of spatial resolutions (from  $0.75^{\circ}$  to  $2.8^{\circ}$ ) and various degrees of sophistication (e.g. biogeochemical cycles, interactive aerosol, cloud-aerosol interaction). Further model details and information on their configuration or features can be found in the PCMDI data portal (<http://www-pcmdi.llnl.gov/>) and in Chapter 9 of the 5<sup>th</sup> IPCC Assessment Report (AR5, IPCC 2013). Model data were first re-gridded to a common  $1^{\circ} \times 1^{\circ}$  reference grid and then averaged to obtain the multi-model mean. Only one ensemble member per model was considered (the *r1i1p1* if available). Although CMIP6 models have just been released, CMIP5 models have been thoroughly tested and broadly used in a vast range of studies and can therefore represent the best well-documented choice for this study (e.g., IPCC 2013).

References are listed in the main text.

### **Text: T3: Metadata concerning highlighted values in Table 1 (main text)**

All values are converted to  $^{\circ}\text{C}/\text{decade}$

Qixiang et al. (2018). IPCC SROCC says  $+0.40/+0.35$  but the original literature calculates  $+0.43$  for the high elevation warming rate which is  $2.165/5$ .

Oyler et al. (2015) The trend given in the abstract, based on the primary analysis and corrections applied in the paper is for 1991-2012 and we use this figure ( $+0.11$ )

Ceppi et al. (2012) The figure given in the abstract is  $+0.48$  (not  $+0.46$  as in SROCC). This is contradicted later in the article so there is a discrepancy.

Yan et al. (2014) Appears to be missing from the final SROCC table, but was included in earlier drafts and so is included here.

**Text: T4: References from meta-study (Table 1 in main text and Tables SM2.2 and SM2.4 in SROCC)**

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		Global	Andes	Tibet	GAR	Rockies	
CRU	1900-2018	1.02/0.95/0.06**	0.34/0.33/0.01	0.79/1.04/-0.24**	1.46/1.44/0.03*	1.05/0.93/0.12	
	1940-2018	1.50/1.50/0.00	0.43/0.52/-0.08	1.16/1.55/-0.39**	2.13/2.15/-0.02	1.49/1.28/0.20	
	1960-2018	2.37/2.27/0.09	0.67/0.63/0.03	2.15/2.32/-0.17	3.36/3.39/-0.03	2.26/2.11/0.16	
	1980-2018	2.75/2.61/0.14	0.47/1.24/-0.78**	2.73/2.81/-0.08	4.51/4.66/-0.15*	2.88/2.20/0.69	
	GISTE MP	1900-2018	1.13/1.13/-0.01	?/1.25/?	1.06/1.11/-0.05	1.28/1.18/0.10**	1.02/0.84/0.19
		1940-2018	1.68/1.68/0.00	1.24/1.30/-0.07	1.42/1.61/-0.20**	1.99/1.88/0.11**	1.69/1.34/0.35
1960-2018		2.59/2.45/0.14	1.53/1.60/-0.07	2.43/2.53/-0.10	3.23/3.08/0.15**	2.59/2.20/0.39	
1980-2018		3.01/2.78/0.23	1.30/1.36/-0.06	3.11/3.16/-0.05	4.54/4.33/0.21**	3.16/2.12/1.04	
ERA5		1980-2018	3.32/3.06/0.27*	2.08/1.94/0.14	3.29/3.00/0.29	5.37/4.44/0.93**	3.66/2.45/1.21*
		1900-2018	0.84/0.83/0.02	0.78/0.67/0.11**	0.73/0.72/0.01	0.72/0.71/0.01*	0.74/0.71/0.03
CMIP5	1940-2018	1.30/1.24/0.06	1.06/0.91/0.15**	1.17/1.15/0.02	1.27/1.28/-0.01	1.24/1.13/0.11*	
	1960-2018	2.42/2.26/0.16**	1.89/1.69/0.20**	2.34/2.27/0.06*	2.55/2.55/0.00	2.48/2.35/0.13	
	1980-2018	3.88/3.42/0.46**	2.38/2.12/0.26**	3.93/3.73/0.19**	4.18/4.13/0.05*	4.27/4.35/-0.08	

**Table S2. Temperature trends (°C/100y) for mountain/lowland/elevation gradient by region for four different periods using four different global gridded datasets (mountain regions defined by boxes in Figure 1). Trend figures show mountain warming rate/lowland warming rate/rate of change in elevation gradient. All absolute trends are positive (warming). Red = mountain warming significantly faster than lowlands. Pink = mountain warming faster (not-significant). Light Orange = mountain warming slower (not significant), Dark Orange = mountain warming significantly slower. Significance of trend in the difference is based on the Mann Kendal test \*\* $p < 0.01$ , \* $p < 0.05$ .**



		Global	Andes	Tibet	GAR	Rockies
CRU	1900-2018	14/22/-8**	4/65/-61**	4/2/2	-20/-21/1	16/42/-26
	1940-2018	8/19/-11*	19/52/-33	4/-20/24	46/16/30	3/47/-44
	1960-2018	1/22/-21**	28/51/-23	24/10/14	-50/-54/4	-15/60/-75
	1980-2018	31/78/-47**	-18/1/-19	45/32/13	74/-7/81	-171/16/-186
	1900-2018	5/11/-7	-3/19/-22	7/19/-12	-4/-12/8	6/19/-13
GPCC	1940-2018	9/18/-9	-16/25/-41	2/9/-6	64/46/18	-6/19/-25
	1960-2018	6/21/-15	-8/19/-27	26/49/-23	-16/-23/7	-22/27/-49
	1980-2018	23/81/-58*	-143/-123/-19	35/32/3	114/74/40	-177/-60/-117
	1980-2018	-96/12/-108**	-337/-182/-155	-223/-66/-157**	103/98/5	-331/-196/-135
	1900-2018	-4/1/-5**	3/2/0	-14/-8/-6**	-16/-10/-5*	-1/2/-3
ERA5	1940-2018	-1/6/-7**	8/-2/10	-11/-4/-7*	13/16/-4	-10/5/-15*
	1960-2018	26/32/-6	32/-3/35*	27/32/-5	61/68/-7	-36/-1/-35**
	1980-2018	70/66/4	29/10/19	91/71/20	136/122/14	-32/17/-48

**Table S3. Trends for mountain precipitation/lowland precipitation/orographic precipitation gradient for individual mountain regions and adjacent lowlands (defined by boxes in Figure 1). Red = significant increased orographic gradient. Pink = increased orographic gradient (not-significant). Light Blue = decreased orographic gradient (not significant). Dark Blue = significant decreased orographic gradient. Significance of trend in the difference is based on the Mann Kendal test \*\* $p < 0.01$ , \* $p < 0.05$ .**